

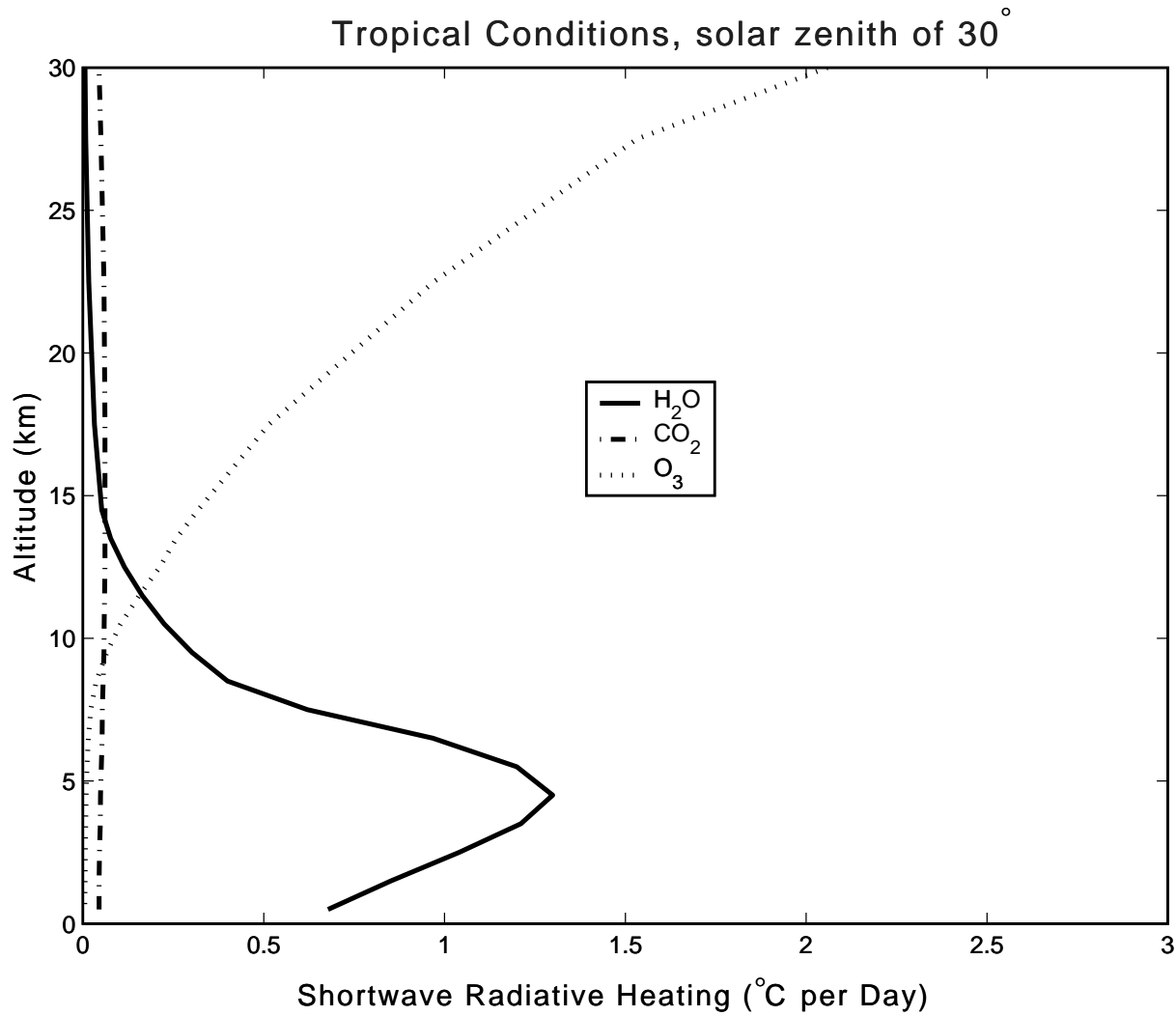
METR 5223: Atmospheric Radiation

Radiative Equilibrium in a Gray Atmosphere

Lecture for Spring 2009, v0.3

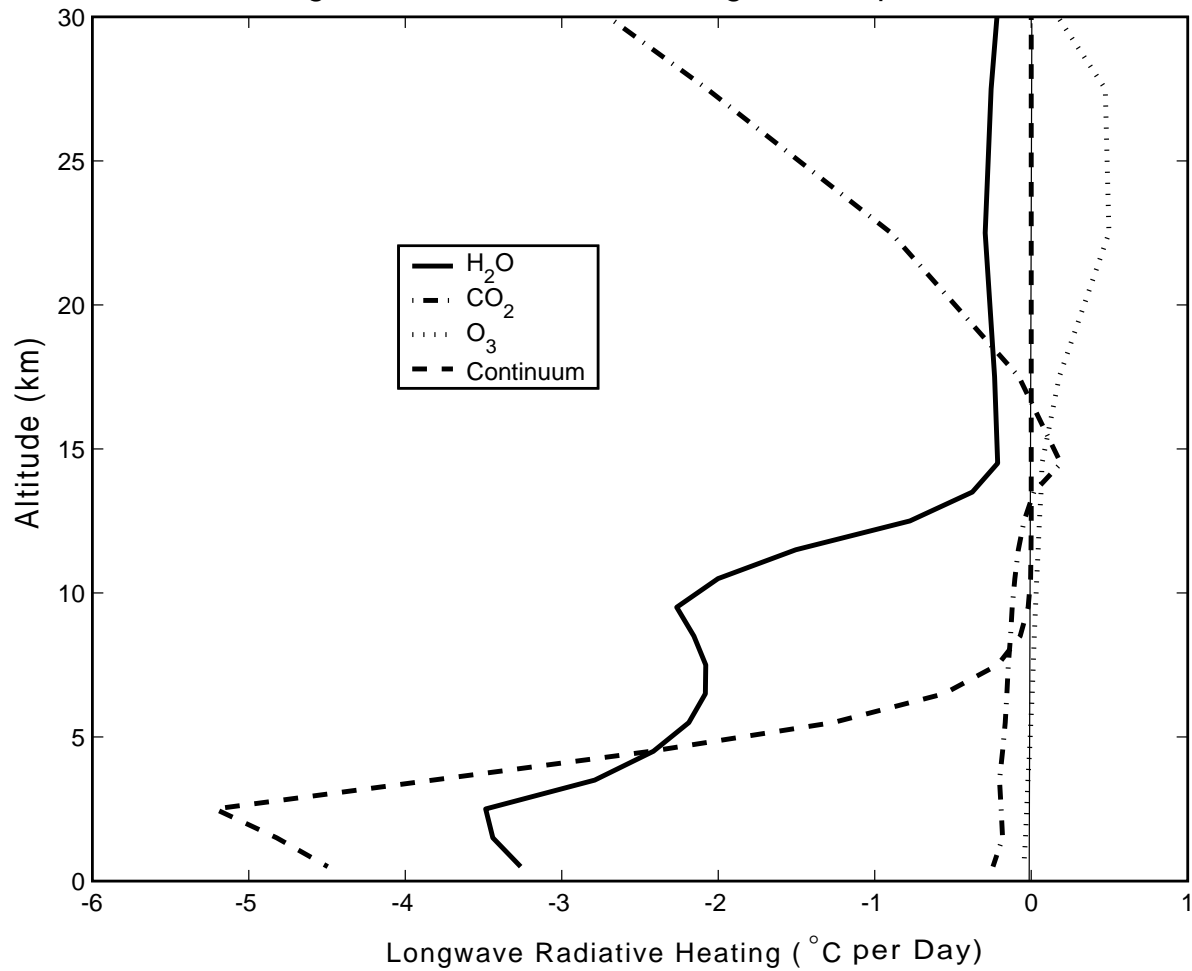
Prof. Brian H. Fiedler

School of Meteorology, University of Oklahoma



Typical heating profiles due to solar absorption in a cloud-free tropical atmosphere.

Longwave Radiative Heating for Tropical Conditions



Typical heating profiles due to longwave radiative transfer in a cloud-free tropical atmosphere.

Idealized radiative equilibrium solutions

- Simplified with all IR having the same value of k_ν .
- No sensible or latent heat fluxes.
- Steady solar heating at surface, and optionally in the atmosphere.
- Simple analytical solution of ODEs.
- Demonstrates stratospheric cooling associated with increased greenhouse gases.

Optical depth at a certain height z is the optical path (thickness) from z to the top of the atmosphere:

$$\chi_\nu(z) = \int_z^\infty k_\nu(z') \rho_a(z') dz'$$

For application to irradiance we use $\chi_\nu^* = 1.66\chi_\nu$ and

$$\frac{dF_\nu \downarrow}{d\chi_\nu^*} = -F_\nu \downarrow + \pi B_\nu$$

Leave the * implied:

$$\frac{dF_{\nu} \downarrow}{d\chi_{\nu}} + F_{\nu} \downarrow = \pi B_{\nu}$$

Also, in a gray atmosphere χ_{ν} and k_{ν} are independent of ν , so $\chi_{\nu} = \chi$. Integrate over ν :

$$\frac{dF \downarrow}{d\chi} + F \downarrow = \sigma T^4$$

Similarly:

$$-\frac{dF \uparrow}{d\chi} + F \uparrow = \sigma T^4$$

$F \downarrow$, $F \uparrow$ and σT^4 are coupled. Need another equation.

In radiative equilibrium, the divergence of the net upward flux is zero:

$$\rho c_p \frac{dT}{dt} = -\frac{d}{dz}(F \uparrow - F \downarrow) = 0$$

So

$$F \uparrow - F \downarrow = F_{\circ}$$

where F_{\circ} is a constant. At $\chi = 0$ (T.O.A), $F \downarrow (0) = 0$ and $F \uparrow (0) = F_{\circ}$. Thus F_{\circ} is also the net downward solar irradiance at the T.O.A..

Recall:

$$\frac{dF \downarrow}{d\chi} + F \downarrow = \sigma T^4$$
$$-\frac{dF \uparrow}{d\chi} + F \uparrow = \sigma T^4$$

Adding the previous 2 Schwarzschild equations together:

$$\frac{d}{d\chi}(F \downarrow - F \uparrow) + F \downarrow + F \uparrow = 2\sigma T^4$$

But

$$\frac{d}{d\chi}(F \downarrow - F \uparrow) = 0$$

So

$$F \downarrow + F \uparrow = 2\sigma T^4$$

$$F \downarrow + F \uparrow = 2\sigma T^4$$

If we multiply both sides by absorptance/emittance:

$$\text{absorption} = \text{emission}$$

Absorption is proportional to the irradiances passing through a layer.

Emission is proportional to $2\sigma T^4$ (the layer radiates in both directions).

Combine

$$F \downarrow + F \uparrow = 2\sigma T^4$$

and

$$F \uparrow - F \downarrow = F_{\circ}$$

to give

$$F \uparrow - F_{\circ} + F \uparrow = 2\sigma T^4$$

or

$$F \uparrow = \sigma T^4 + \frac{1}{2}F_{\circ}$$

and

$$F \downarrow = \sigma T^4 - \frac{1}{2}F_{\circ}$$

Rewrite previous equation as:

$$\sigma T^4 = F \downarrow + \frac{1}{2} F_{\circ}$$

Substitute into

$$\frac{dF \downarrow}{d\chi} + F \downarrow = \sigma T^4$$

$$\frac{dF \downarrow}{d\chi} + F \downarrow = F \downarrow + \frac{1}{2} F_{\circ}$$

Solution is simple. Using $F \downarrow (0) = 0$,

$$F \downarrow (\chi) = \frac{1}{2} F_{\circ} \chi$$

Recap:

$$F \downarrow (\chi) = \frac{1}{2} F_{\circ} \chi$$

$$\sigma T^4(\chi) = \frac{1}{2} F_{\circ} \chi + \frac{1}{2} F_{\circ}$$

$$F \uparrow (\chi) = \frac{1}{2} F_{\circ} \chi + F_{\circ}$$

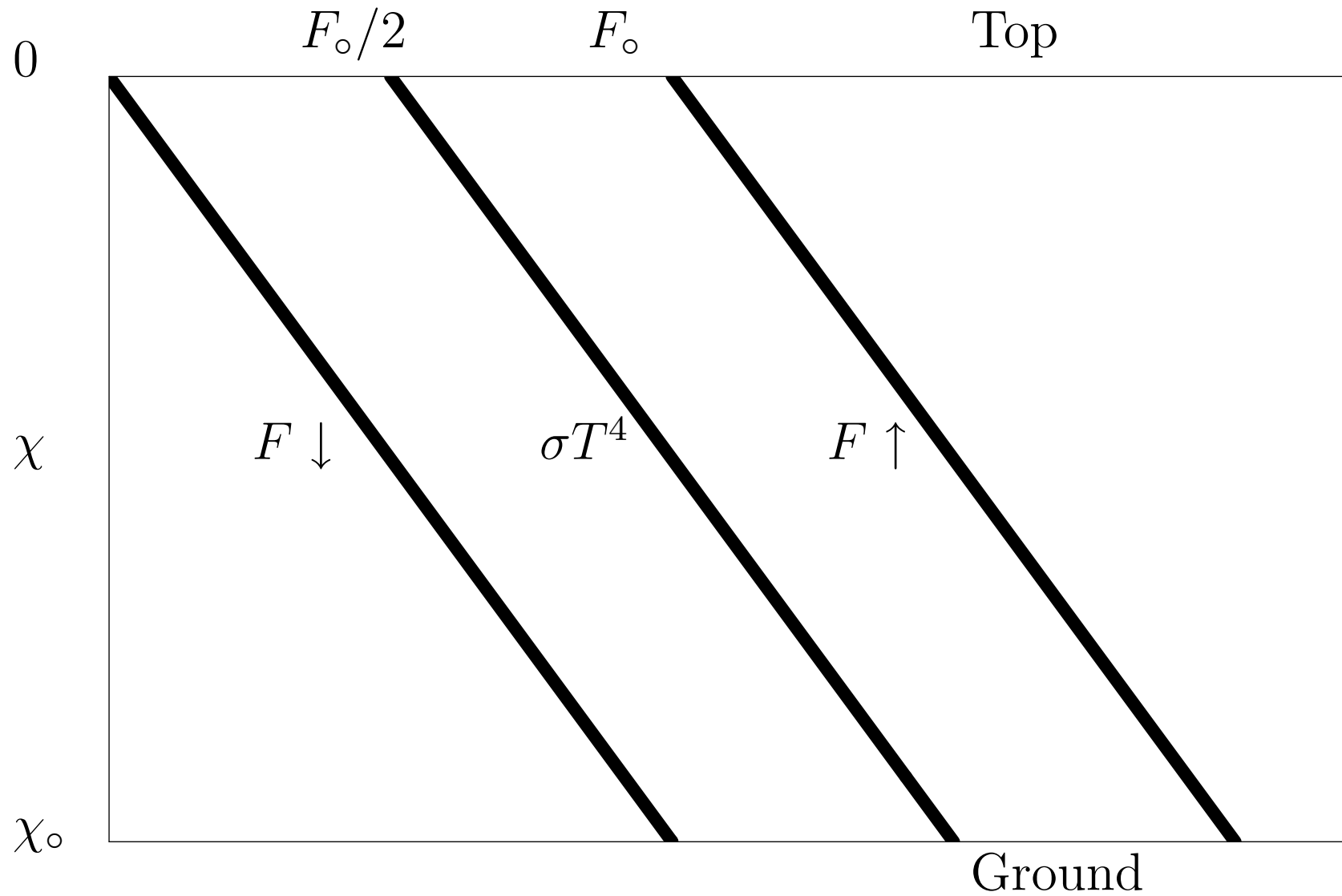
Note at the ground, or $\chi = \chi_{\circ}$, T_g is given by:

$$F \uparrow (\chi_{\circ}) = \frac{1}{2} F_{\circ} \chi_{\circ} + F_{\circ} = \sigma T_g^4$$

But the air temperature at the ground is T_a :

$$\sigma T_a^4 = \sigma T_g^4 - \frac{1}{2} F_{\circ}$$

Radiative Equilibrium in a Grey Atmosphere



Generalize the gray atmosphere model to include a prescribed solar absorption acting on a downward solar irradiance Q .

Model is more clear using dimensionless pressure p for the independent variable. $p = 0$ at the top of the atmosphere and $p = 1$ at the surface.

The optical path (optical depth) for longwave radiation is now simply $\chi = \chi_0 p$.

$$Q = F_o - sF_o (1 - e^{-\alpha p})$$

α controls the rate of reduction with p .

s controls the fraction of Q that is ultimately removed.

We do experiments for increasing χ_o (the value of χ at the surface) to explore an increased greenhouse effect.

We will hold α and s invariant, which keeps solar absorption invariant in its magnitude and position.

The solar irradiance Q is positive downward.

The solar heating is proportional to the convergence of the solar irradiance, or

$$-\frac{dQ}{dp}$$

The longwave heating is proportional to the convergence of the net downward irradiance,

$$-\frac{d}{dp} (F \downarrow - F \uparrow)$$

There is no heating by latent heat release or convection in the gray atmosphere model. The sum of the radiational heating is zero:

$$-\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

In equilibrium, there is 0 net irradiance at the top of the atmosphere. So the above equation integrates to

$$Q + F \downarrow - F \uparrow = 0$$

for all p .

Recall the Schwarzschild equations for the downward and upward longwave irradiance:

$$\frac{1}{\chi_o} \frac{dF \downarrow}{dp} + F \downarrow = \sigma T^4$$

$$-\frac{1}{\chi_o} \frac{dF \uparrow}{dp} + F \uparrow = \sigma T^4$$

Recall the goal here is to solve for $F \uparrow$, $F \downarrow$ and σT^4 using the boundary conditions $F \uparrow(0) = F_o$ and $F \downarrow(0) = 0$, and the zero net heating condition:

$$\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

$$\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

can be written

$$\frac{dQ}{dp} = \frac{dF \uparrow}{dp} - \frac{dF \downarrow}{dp}$$

$$\frac{dQ}{dp} = \chi_o (F \downarrow + F \uparrow - 2\sigma T^4)$$

$$\frac{dQ}{dp} = \chi_o (2F \downarrow + Q - 2\sigma T^4)$$

Using

$$\frac{dQ}{dp} = \chi_o (2F \downarrow + Q - 2\sigma T^4)$$

to eliminate σT^4 from

$$\frac{1}{\chi_o} \frac{dF \downarrow}{dp} + F \downarrow = \sigma T^4$$

gives

$$\frac{1}{\chi_o} \frac{d}{dp} \left(F \downarrow + \frac{1}{2} Q \right) = \frac{1}{2} Q$$

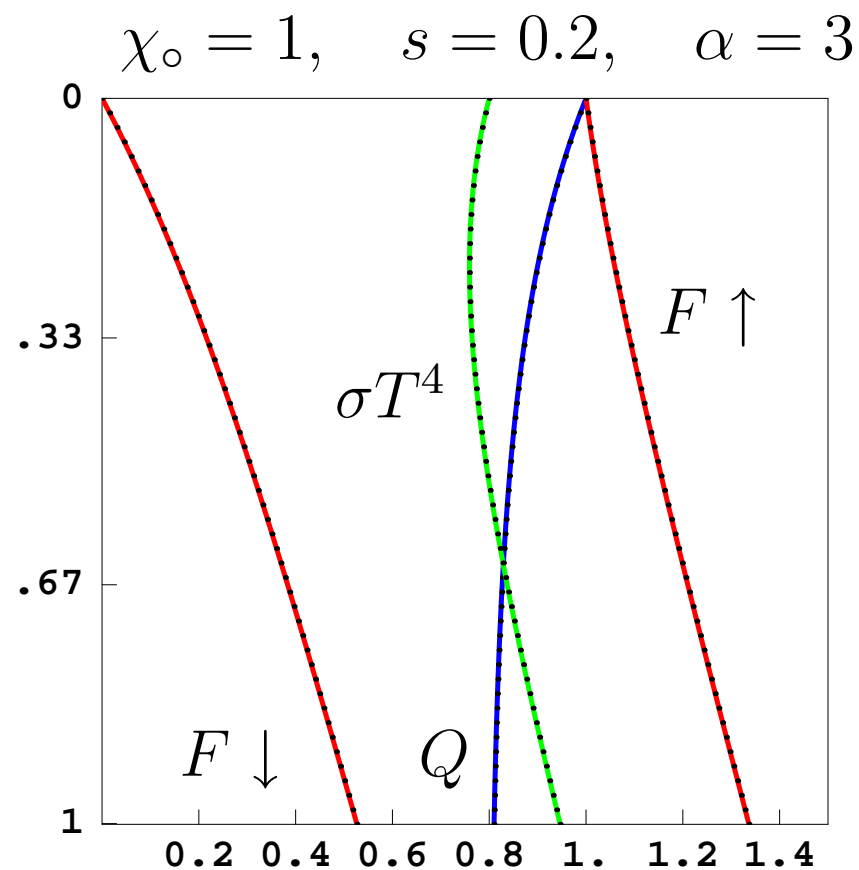
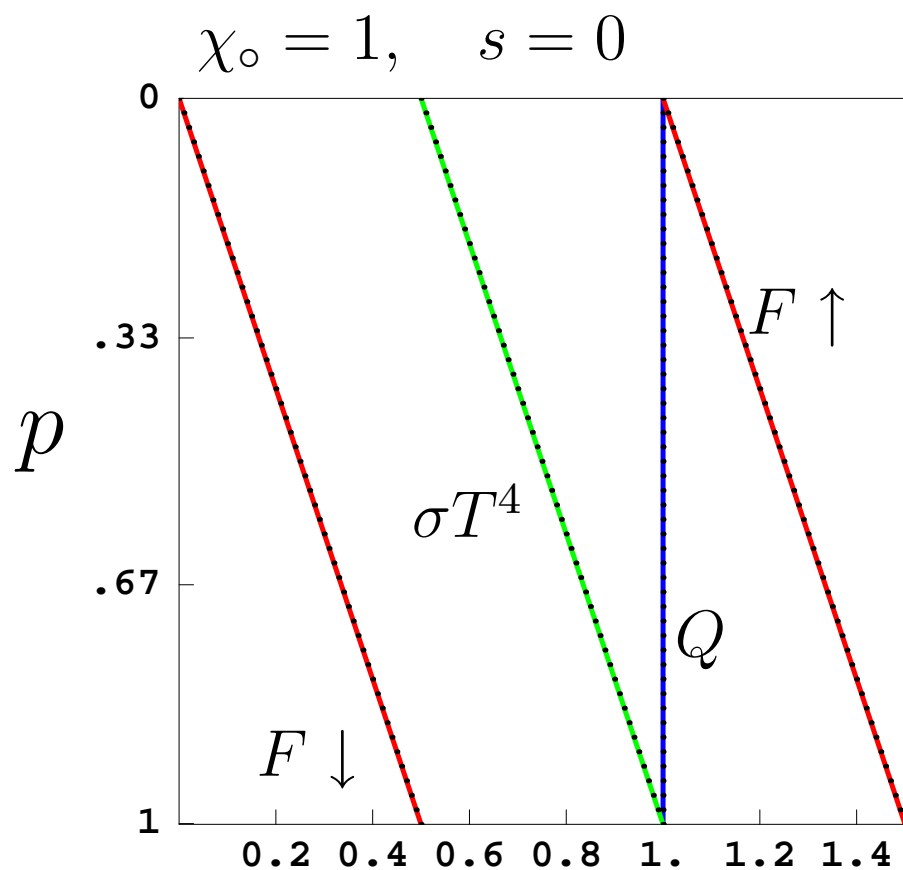
This has solution

$$F \downarrow = \frac{1}{2} F_o (1 - s) \chi_o p + \frac{1}{2} F_o s \left(1 + \frac{\chi_o}{\alpha} \right) (1 - e^{-\alpha p})$$

Using the solution for $F \downarrow (p)$, we obtain $\sigma T^4(p)$ and $F \uparrow (p)$ using:

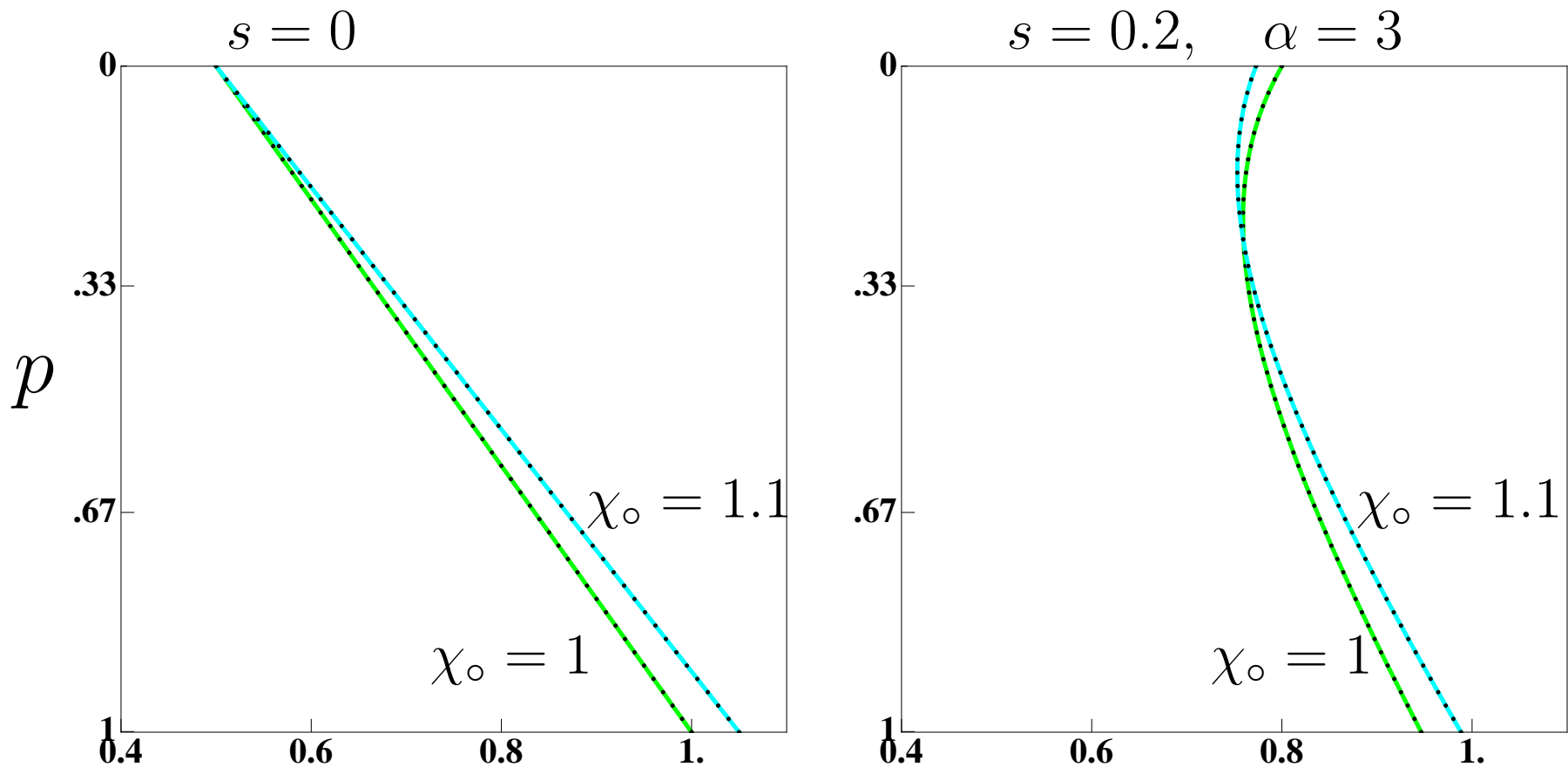
$$\sigma T^4 = F \downarrow + \frac{1}{\chi_o} \frac{dF \downarrow}{dp}$$

$$F \uparrow = F \downarrow + Q$$



Gray-atmosphere solution. Left: no solar absorption. Right: with solar absorption. All flux values normalized by F_{\circ} .

As χ_o increases, the surface air warms but the stratosphere cools:



σT^4 (normalized by F_o) for $\chi_o = 1$ and $\chi_o = 1.1$

Why? Recall the zero net heating condition:

$$-\frac{dQ}{dp} + \frac{dF \uparrow}{dp} - \frac{dF \downarrow}{dq} = 0$$

$$-\frac{dQ}{dp} + \chi_{\circ} (F \downarrow + F \uparrow - 2\sigma T^4) = 0$$

Applied to the top of the atmosphere:

$$\alpha F_{\circ} s + \chi_{\circ} (F_{\circ} - 2\sigma T^4) = 0$$

$\alpha F_{\circ} s$ is solar heating. $\chi_{\circ} F_{\circ}$ is longwave heating. $\chi_{\circ} 2\sigma T^4$ is longwave cooling.

This can be rearranged as:

$$\frac{1}{F_{\circ}} \sigma T^4 = \frac{1}{2} + \frac{\alpha}{2\chi_{\circ}} s$$

As χ_{\circ} increases, σT^4 will decrease.

This allows longwave cooling to balance the solar heating, which is assumed invariant here.