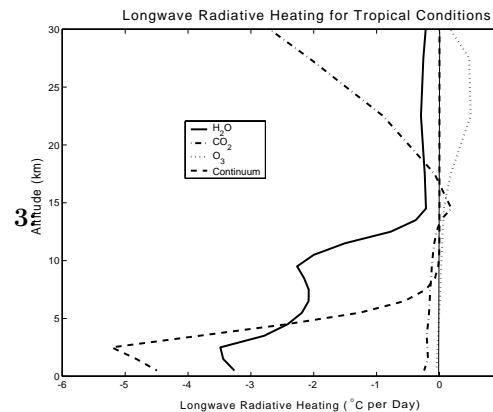


Radiative Equilibrium in a Gray Atmosphere

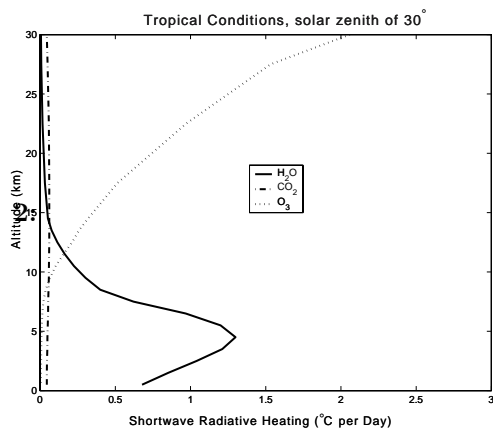
1:

Lecture for Spring 2009, v0.3
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Typical heating profiles due to longwave radiative transfer in a cloud-free tropical atmosphere.



Typical heating profiles due to solar absorption in a cloud-free tropical atmosphere.

- 4:
- Idealized radiative equilibrium solutions**
- Simplified with all IR having the same value of k_ν .
 - No sensible or latent heat fluxes.
 - Steady solar heating at surface, and optionally in the atmosphere.
 - Simple analytical solution of ODEs.
 - Demonstrates stratospheric cooling associated with increased greenhouse gases.

Optical depth at a certain height z is the optical path (thickness) from z to the top of the atmosphere:

$$\chi_\nu(z) = \int_z^\infty k_\nu(z') \rho_a(z') dz'$$

5:

For application to irradiance we use $\chi_\nu^* = 1.66\chi_\nu$ and

$$\frac{dF_\nu \downarrow}{d\chi_\nu^*} = -F_\nu \downarrow + \pi B_\nu$$

Leave the * implied:

$$\frac{dF_\nu \downarrow}{d\chi_\nu} + F_\nu \downarrow = \pi B_\nu$$

Also, in a **gray atmosphere** χ_ν and k_ν are independent of ν , so $\chi_\nu = \chi$. Integrate over ν :

6:

$$\frac{dF \downarrow}{d\chi} + F \downarrow = \sigma T^4$$

Similarly:

$$-\frac{dF \uparrow}{d\chi} + F \uparrow = \sigma T^4$$

$F \downarrow$, $F \uparrow$ and σT^4 are coupled. Need another equation.

In radiative equilibrium, the divergence of the net upward flux is zero:

$$\rho c_p \frac{dT}{dt} = -\frac{d}{dz}(F \uparrow - F \downarrow) = 0$$

7: So

$$F \uparrow - F \downarrow = F_o$$

where F_o is a constant. At $\chi = 0$ (T.O.A), $F \downarrow(0) = 0$ and $F \uparrow(0) = F_o$. Thus F_o is also the net downward solar irradiance at the T.O.A..

Recall:

$$\begin{aligned} \frac{dF \downarrow}{d\chi} + F \downarrow &= \sigma T^4 \\ -\frac{dF \uparrow}{d\chi} + F \uparrow &= \sigma T^4 \end{aligned}$$

Adding the previous 2 Schwarzschild equations together:

8:

$$\frac{d}{d\chi}(F \downarrow - F \uparrow) + F \downarrow + F \uparrow = 2\sigma T^4$$

But

$$\frac{d}{d\chi}(F \downarrow - F \uparrow) = 0$$

So

$$F \downarrow + F \uparrow = 2\sigma T^4$$

$$F \downarrow + F \uparrow = 2\sigma T^4$$

If we multiply both sides by absorptance/emittance:

$$\text{absorption} = \text{emission}$$

9:

Absorption is proportional to the irradiances passing through a layer.

Emission is proportional to $2\sigma T^4$ (the layer radiates in both directions).

Combine

$$F \downarrow + F \uparrow = 2\sigma T^4$$

and

$$F \uparrow - F \downarrow = F_0$$

to give

10:

$$F \uparrow - F_0 + F \uparrow = 2\sigma T^4$$

or

$$F \uparrow = \sigma T^4 + \frac{1}{2}F_0$$

and

$$F \downarrow = \sigma T^4 - \frac{1}{2}F_0$$

Rewrite previous equation as:

$$\sigma T^4 = F \downarrow + \frac{1}{2}F_0$$

Substitute into

$$\frac{dF \downarrow}{d\chi} + F \downarrow = \sigma T^4$$

11:

$$\frac{dF \downarrow}{d\chi} + F \downarrow = F \downarrow + \frac{1}{2}F_0$$

Solution is simple. Using $F \downarrow(0) = 0$,

$$F \downarrow(\chi) = \frac{1}{2}F_0\chi$$

Recap:

$$F \downarrow(\chi) = \frac{1}{2}F_0\chi$$

$$\sigma T^4(\chi) = \frac{1}{2}F_0\chi + \frac{1}{2}F_0$$

$$F \uparrow(\chi) = \frac{1}{2}F_0\chi + F_0$$

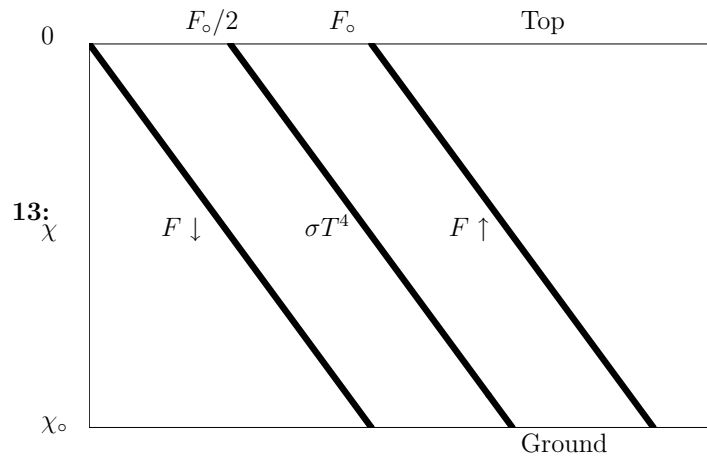
12: Note at the ground, or $\chi = \chi_0$, T_g is given by:

$$F \uparrow(\chi_0) = \frac{1}{2}F_0\chi_0 + F_0 = \sigma T_g^4$$

But the air temperature at the ground is T_a :

$$\sigma T_a^4 = \sigma T_g^4 - \frac{1}{2}F_0$$

Radiative Equilibrium in a Grey Atmosphere



13: χ

Generalize the gray atmosphere model to include a prescribed solar absorption acting on a downward solar irradiance Q .

14:

Model is more clear using dimensionless pressure p for the independent variable. $p = 0$ at the top of the atmosphere and $p = 1$ at the surface.

The optical path (optical depth) for longwave radiation is now simply $\chi = \chi_0 p$.

$$Q = F_o - sF_o (1 - e^{-\alpha p})$$

15:

α controls the rate of reduction with p .
 s controls the fraction of Q that is ultimately removed.

16:

We do experiments for increasing χ_0 (the value of χ at the surface) to explore an increased greenhouse effect.
 We will hold α and s invariant, which keeps solar absorption invariant in its magnitude and position.

The solar irradiance Q is positive downward.

The solar heating is proportional to the convergence of the solar irradiance, or

17:
$$-\frac{dQ}{dp}$$

The longwave heating is proportional to the convergence of the net downward irradiance,

$$-\frac{d}{dp} (F \downarrow - F \uparrow)$$

There is no heating by latent heat release or convection in the gray atmosphere model. The sum of the radiational heating is zero:

$$-\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

18: In equilibrium, there is 0 net irradiance at the top of the atmosphere. So the above equation integrates to

$$Q + F \downarrow - F \uparrow = 0$$

for all p .

Recall the Schwarzschild equations for the downward and upward longwave irradiance:

$$\frac{1}{\chi_o} \frac{dF \downarrow}{dp} + F \downarrow = \sigma T^4$$

$$-\frac{1}{\chi_o} \frac{dF \uparrow}{dp} + F \uparrow = \sigma T^4$$

19:

Recall the goal here is to solve for $F \uparrow$, $F \downarrow$ and σT^4 using the boundary conditions $F \uparrow(0) = F_o$ and $F \downarrow(0) = 0$, and the zero net heating condition:

$$\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

$$\frac{d}{dp} (Q + F \downarrow - F \uparrow) = 0$$

can be written

$$\frac{dQ}{dp} = \frac{dF \uparrow}{dp} - \frac{dF \downarrow}{dp}$$

20:

$$\frac{dQ}{dp} = \chi_o (F \downarrow + F \uparrow - 2\sigma T^4)$$

$$\frac{dQ}{dp} = \chi_o (2F \downarrow + Q - 2\sigma T^4)$$

Using

$$\frac{dQ}{dp} = \chi_o (2F \downarrow + Q - 2\sigma T^4)$$

to eliminate σT^4 from

$$\frac{1}{\chi_o} \frac{dF \downarrow}{dp} + F \downarrow = \sigma T^4$$

21: gives

$$\frac{1}{\chi_o} \frac{d}{dp} \left(F \downarrow + \frac{1}{2} Q \right) = \frac{1}{2} Q$$

This has solution

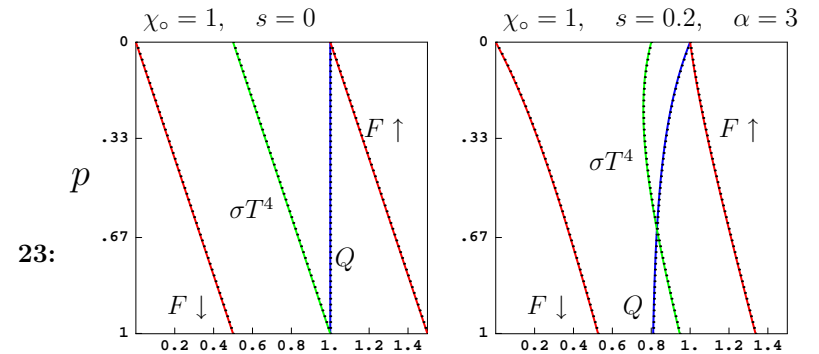
$$F \downarrow = \frac{1}{2} F_o (1 - s) \chi_o p + \frac{1}{2} F_o s \left(1 + \frac{\chi_o}{\alpha} \right) (1 - e^{-\alpha p})$$

Using the solution for $F \downarrow(p)$, we obtain $\sigma T^4(p)$ and $F \uparrow(p)$ using:

22:

$$\sigma T^4 = F \downarrow + \frac{1}{\chi_o} \frac{dF \downarrow}{dp}$$

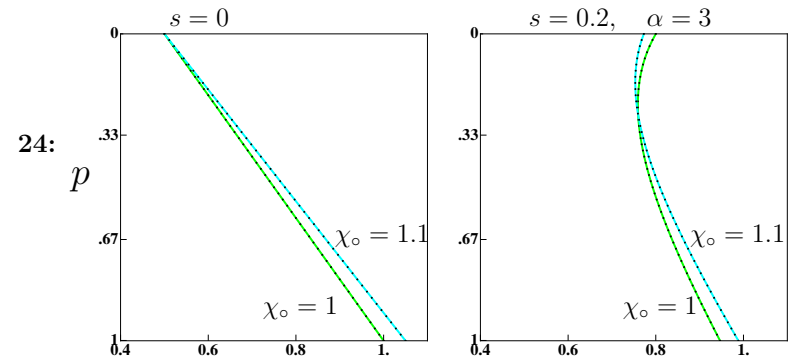
$$F \uparrow = F \downarrow + Q$$



Gray-atmosphere solution. Left: no solar absorption.

Right: with solar absorption. All flux values normalized by F_o .

As χ_o increases, the surface air warms but the stratosphere cools:



σT^4 (normalized by F_o) for $\chi_o = 1$ and $\chi_o = 1.1$

Why? Recall the zero net heating condition:

$$-\frac{dQ}{dp} + \frac{dF \uparrow}{dp} - \frac{dF \downarrow}{dq} = 0$$

$$-\frac{dQ}{dp} + \chi_o (F \downarrow + F \uparrow - 2\sigma T^4) = 0$$

25:

Applied to the top of the atmosphere:

$$\alpha F_o s + \chi_o (F_o - 2\sigma T^4) = 0$$

$\alpha F_o s$ is solar heating. $\chi_o F_o$ is longwave heating. $\chi_o 2\sigma T^4$ is longwave cooling.

This can be rearranged as:

$$\frac{1}{F_o} \sigma T^4 = \frac{1}{2} + \frac{\alpha}{2\chi_o} s$$

26:

As χ_o increases, σT^4 will decrease.

This allows longwave cooling to balance the solar heating, which is assumed invariant here.